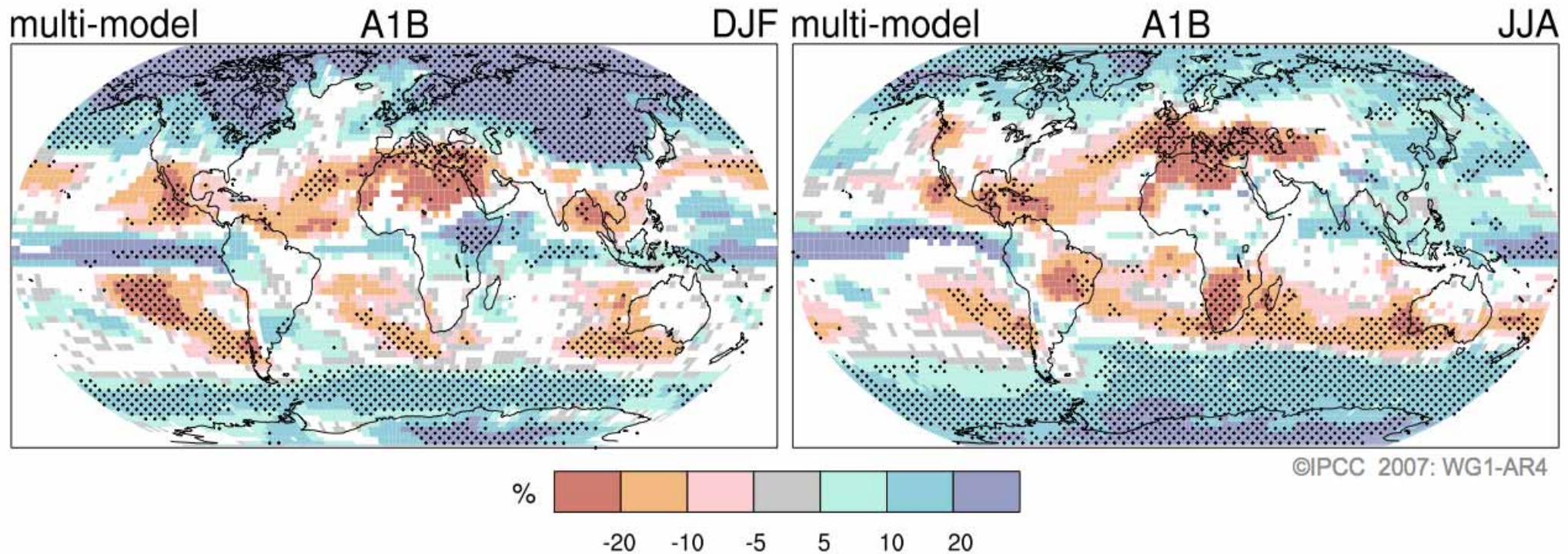


# **Water Cycle Projection in Asia by Super-High-Resolution Climate Model**

**Akio Kitoh**

**Meteorological Research Institute**

# Precipitation changes

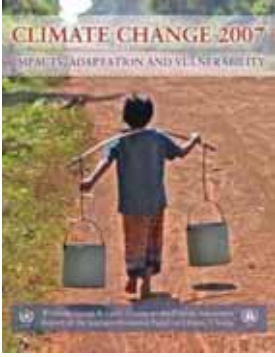


White: <2/3 of models agree on sign of change

Stippled: >90% of models agree on sign of change

Precipitation **increases** *very likely* in high latitudes

**Decreases** *likely* in most subtropical land regions



# Impacts: Freshwater

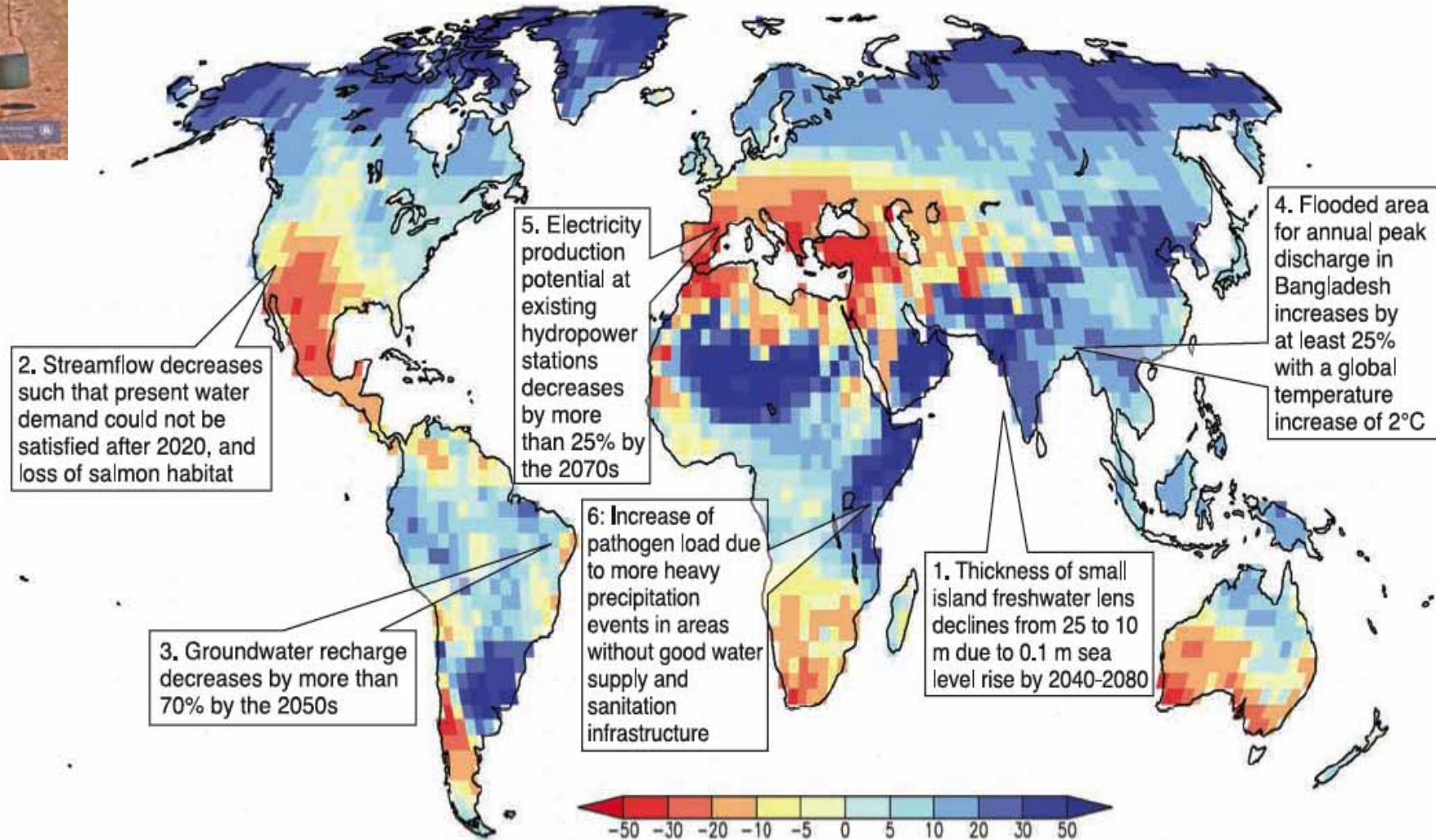
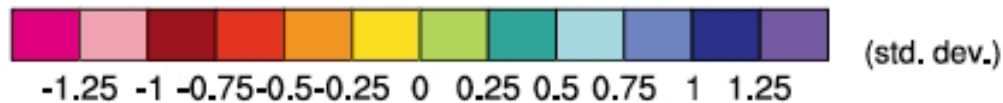
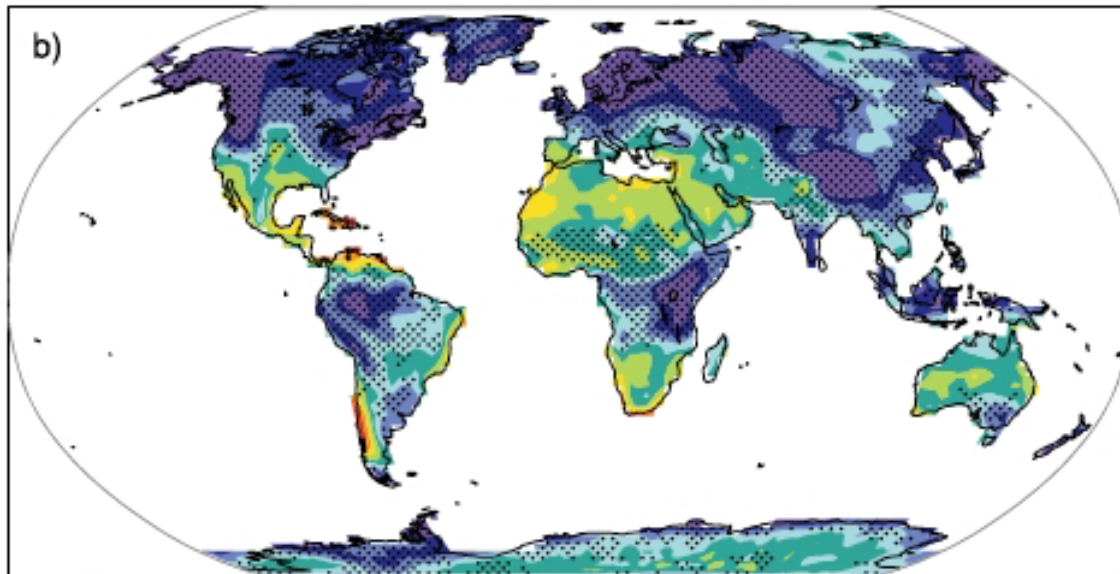


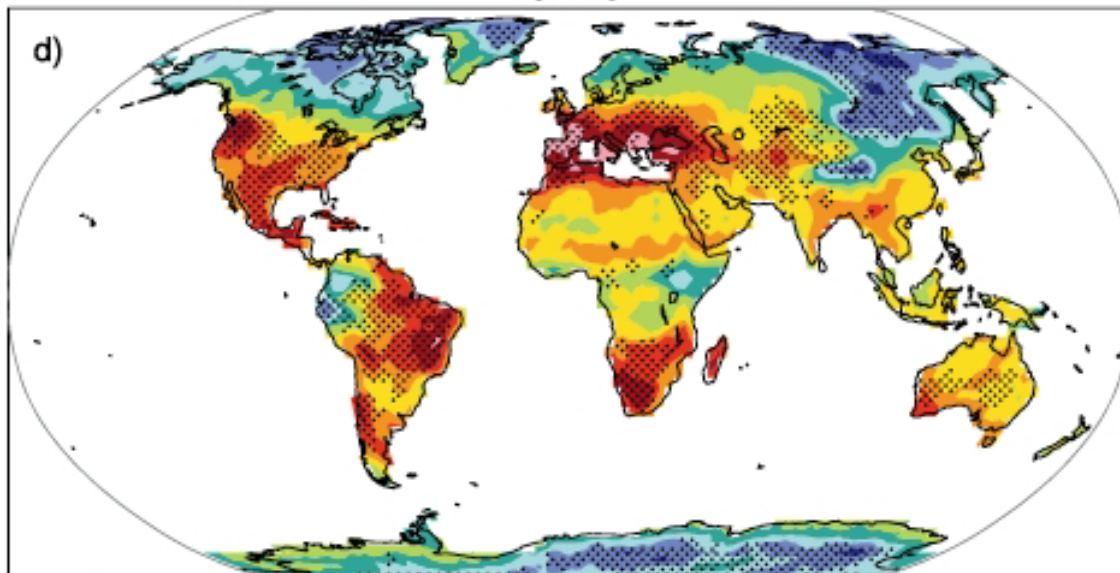
Figure 3.8. Illustrative map of future climate change impacts on freshwater which are a threat to the sustainable development of the affected regions. 1: Bobba et al. (2000), 2: Barnett et al. (2004), 3: Döll and Flörke (2005), 4: Mirza et al. (2003) 5: Lehner et al. (2005a) 6: Kistemann et al. (2002). Background map: Ensemble mean change of annual runoff, in percent, between present (1981 to 2000) and 2081 to 2100 for the SRES A1B emissions scenario (after Nohara et al., 2006).

Increases in the frequency of droughts and floods are projected to affect local crop production negatively, especially in subsistence sectors at low latitudes.

Precipitation intensity



Dry days



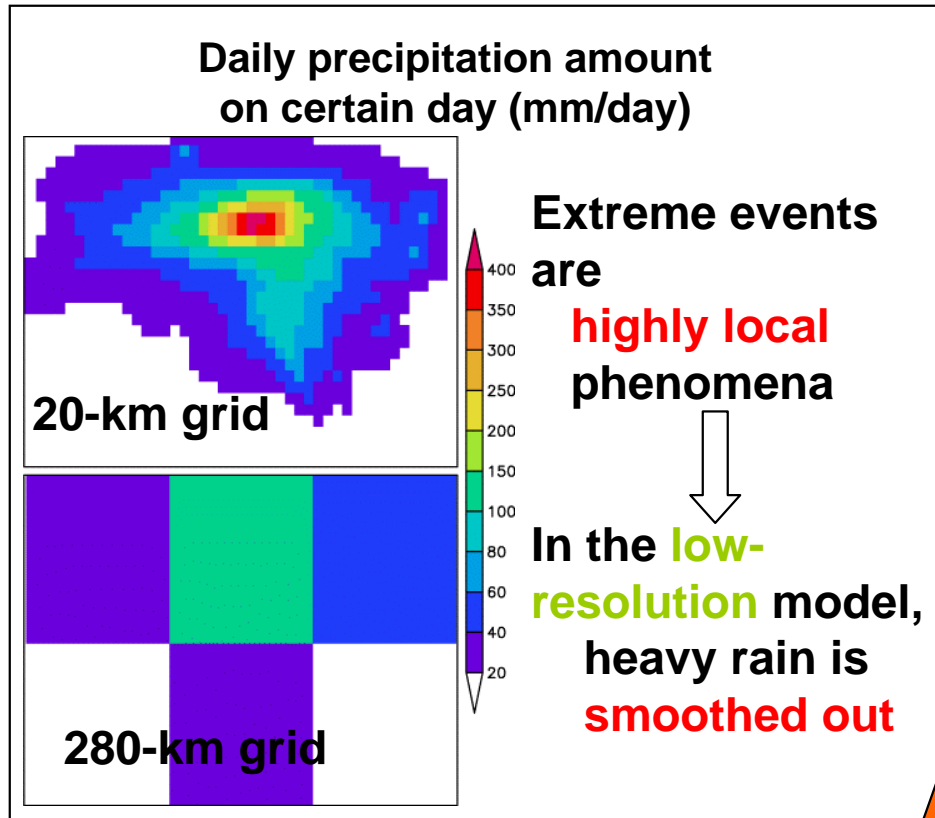
## Projected changes in extremes

Warming of day and night extreme temperatures is *virtually certain*.

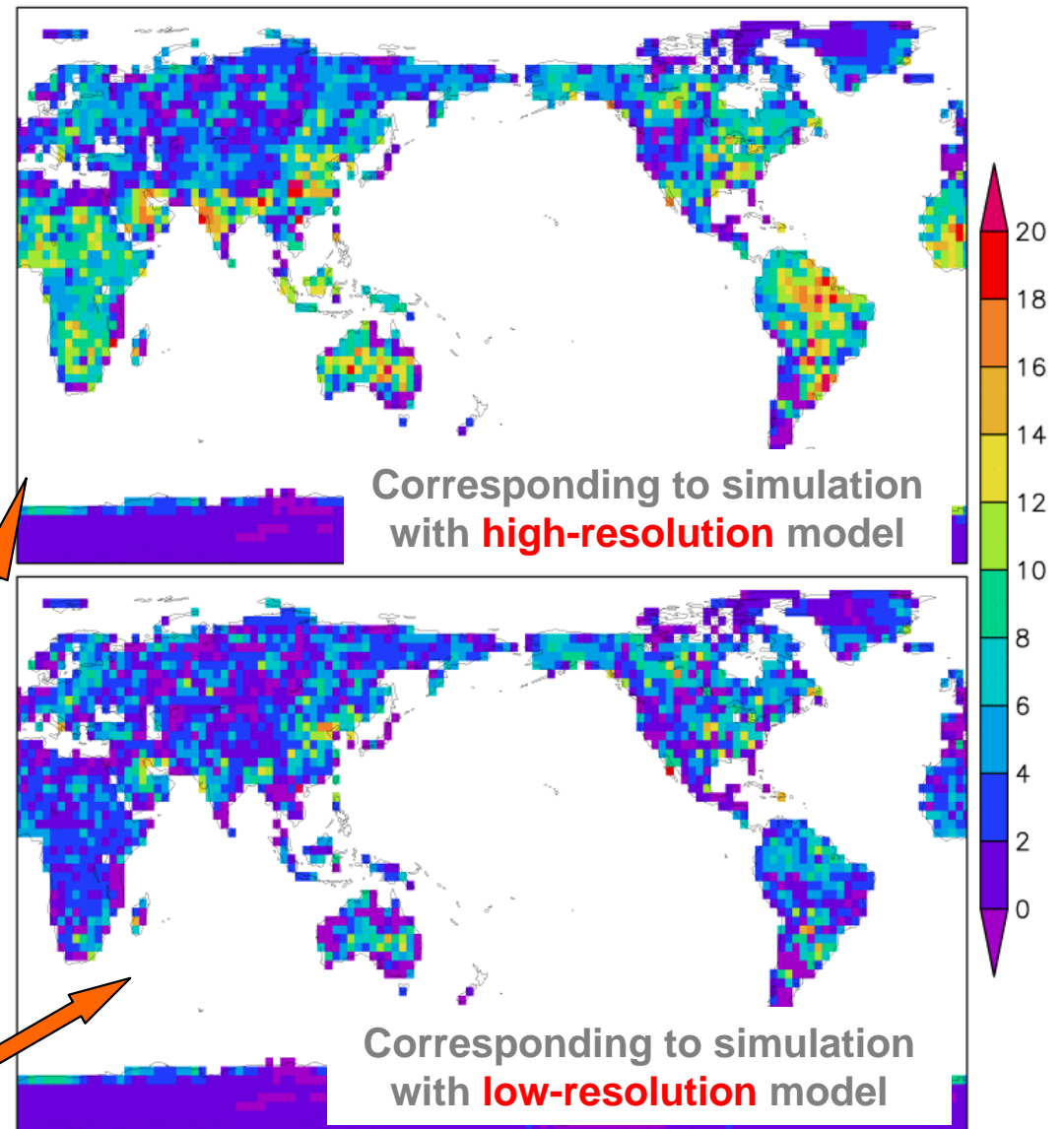
It is *very likely* that heat waves and heavy precipitation events will continue to become more frequent.

Based on a range of models, it is *likely* that future tropical cyclones will become more intense.

# Needs for High-Resolution Model



Future Change in the Annual Maximum Precipitation  
Future – Present (mm/day)



Right upper

The annual maximum rain calculated by original 20-km mesh daily rain

To compare with the lower figure, the annual maximum rain is regridded to low resolution (280-km) grid

Right lower

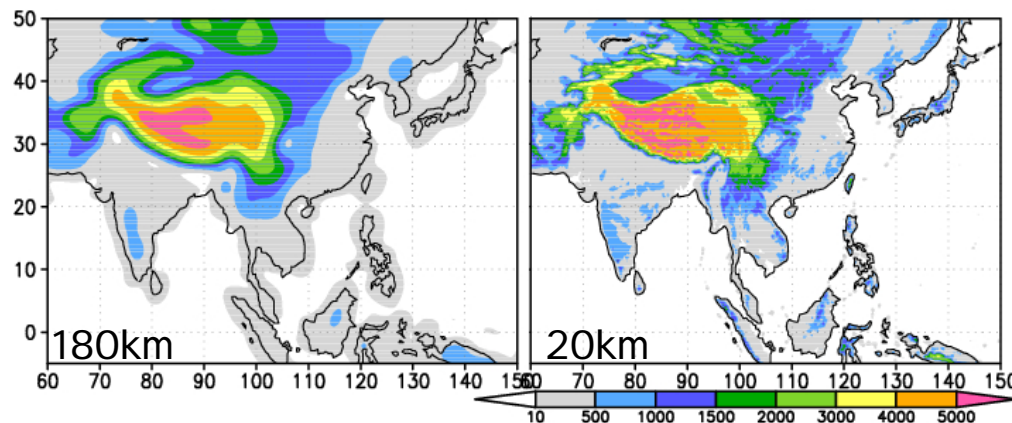
The annual maximum rain calculated by regridded daily rain, which is interpolated to 280-km grid from original 20-km grid

This can be considered as a proxy of low resolution model

**Change in precipitation**

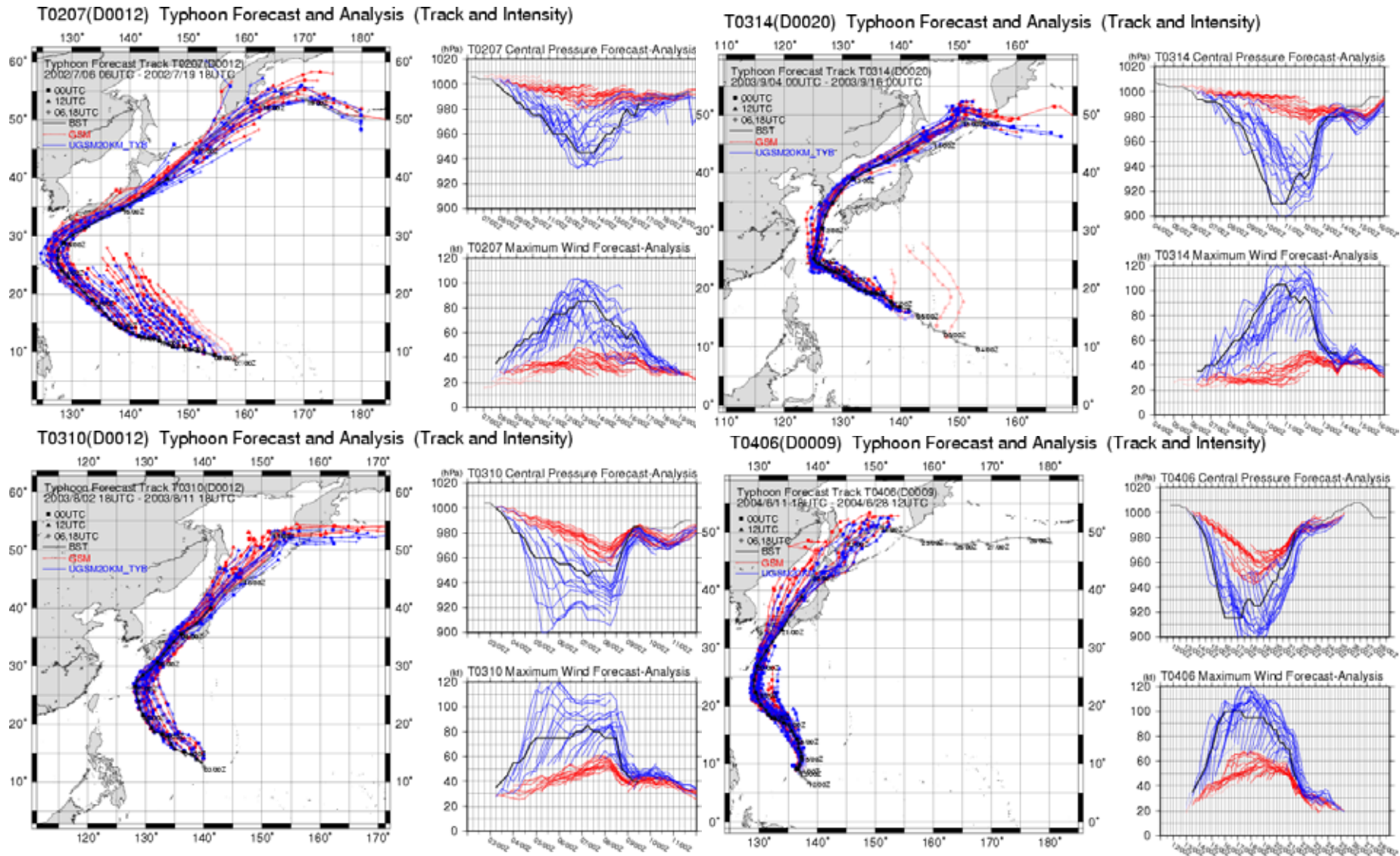
# MRI / JMA Atmospheric GCM

- JMA : Operational global NWP model from Nov 2007
- MRI : Next generation climate model
- Based on **operational JMA-GSM**
- Resolution: **TL959(20km) with 60 layers**
- Time integration: **Semi-Lagrangian Scheme** (Yoshimura, 2004)  
**2 days/1 year integration with DT=6 min and  
30 nodes of Earth Simulator (ES has total 640 nodes)**
- Physics
  - SW radiation: Shibata & Uchiyama (1992)
  - LW radiation: Shibata & Aoki (1989)
  - Cumulus convection: Prognostic Arakawa-Schubert (Randall and Pan, 1993)
  - Land hydrology: MJ-SiB: SiB with 4 soil-layers and 3 snow-layers
  - Clouds: large-scale condensation, Cumulus, stratocumulus
  - PBL: Mellor & Yamada (1974,1982) level-2 closure model
  - Gravity wave drag: Iwasaki et al. (1989) + Rayleigh friction

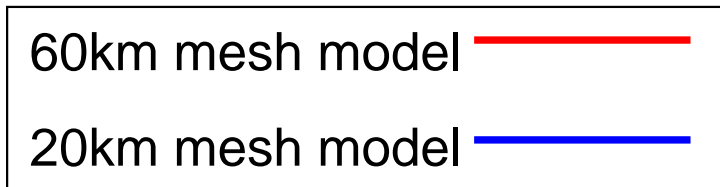


This model is used in MRI -CGCM3 after introducing additional physics and tuning.

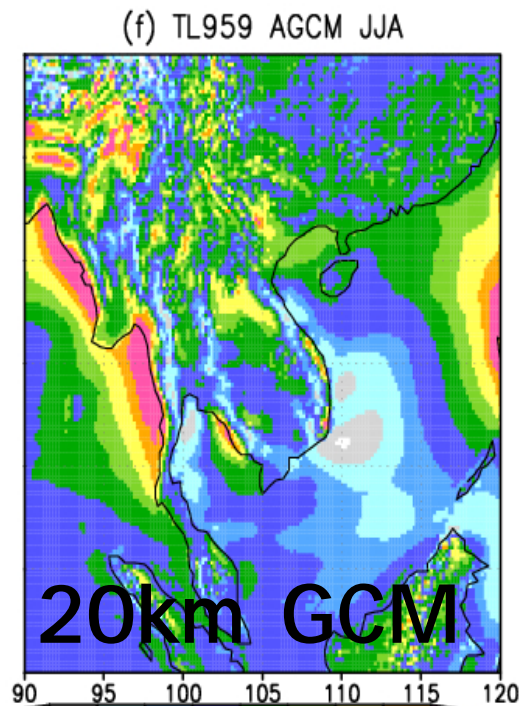
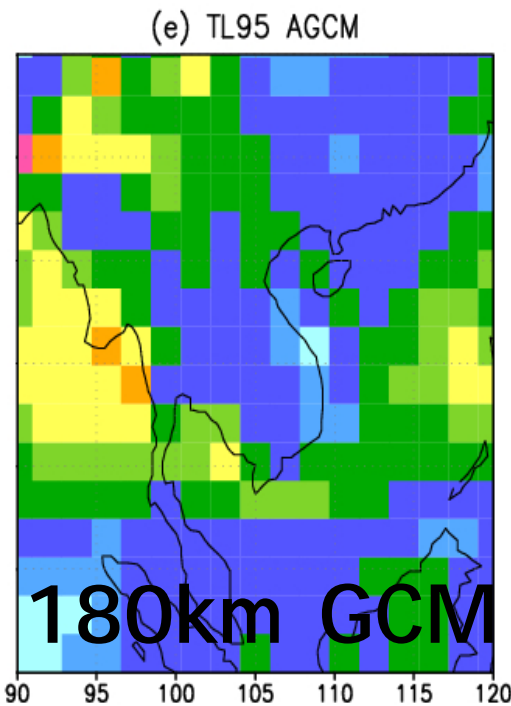
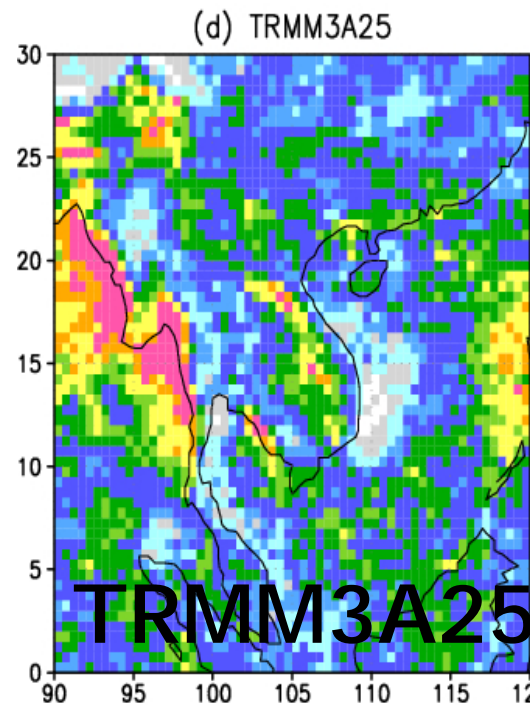
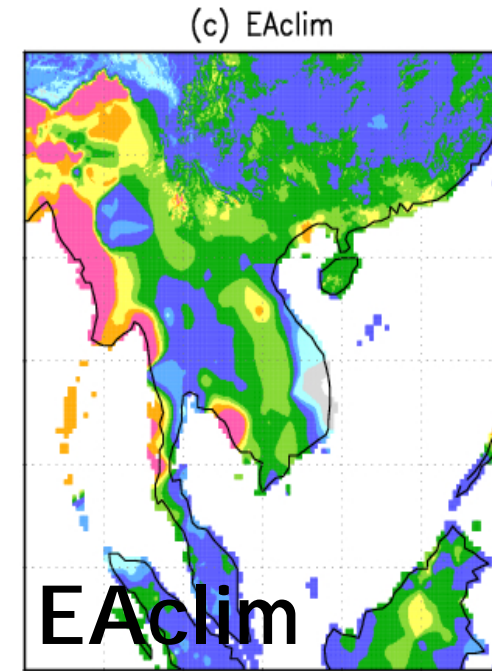
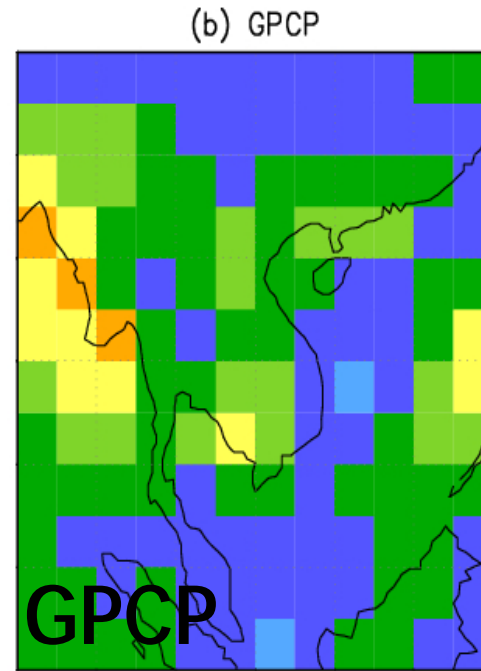
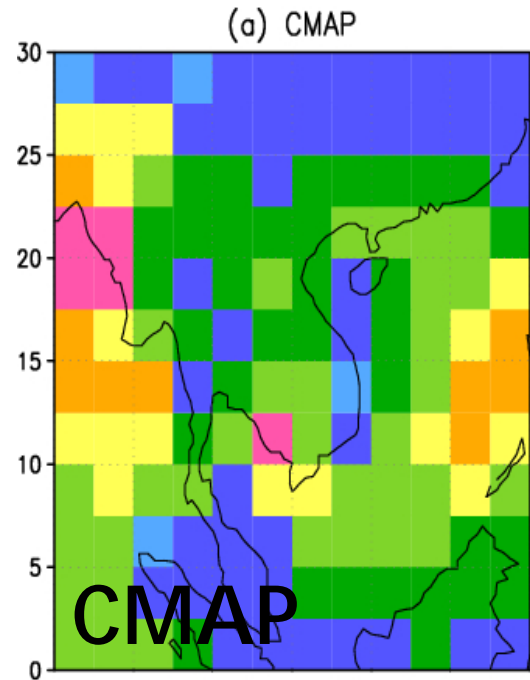
# Typhoon track and intensity: 60km vs 20km



60-km model forecasts shallower central pressures and weaker maximum winds. 20-km model represents typhoon development closer to the observations.

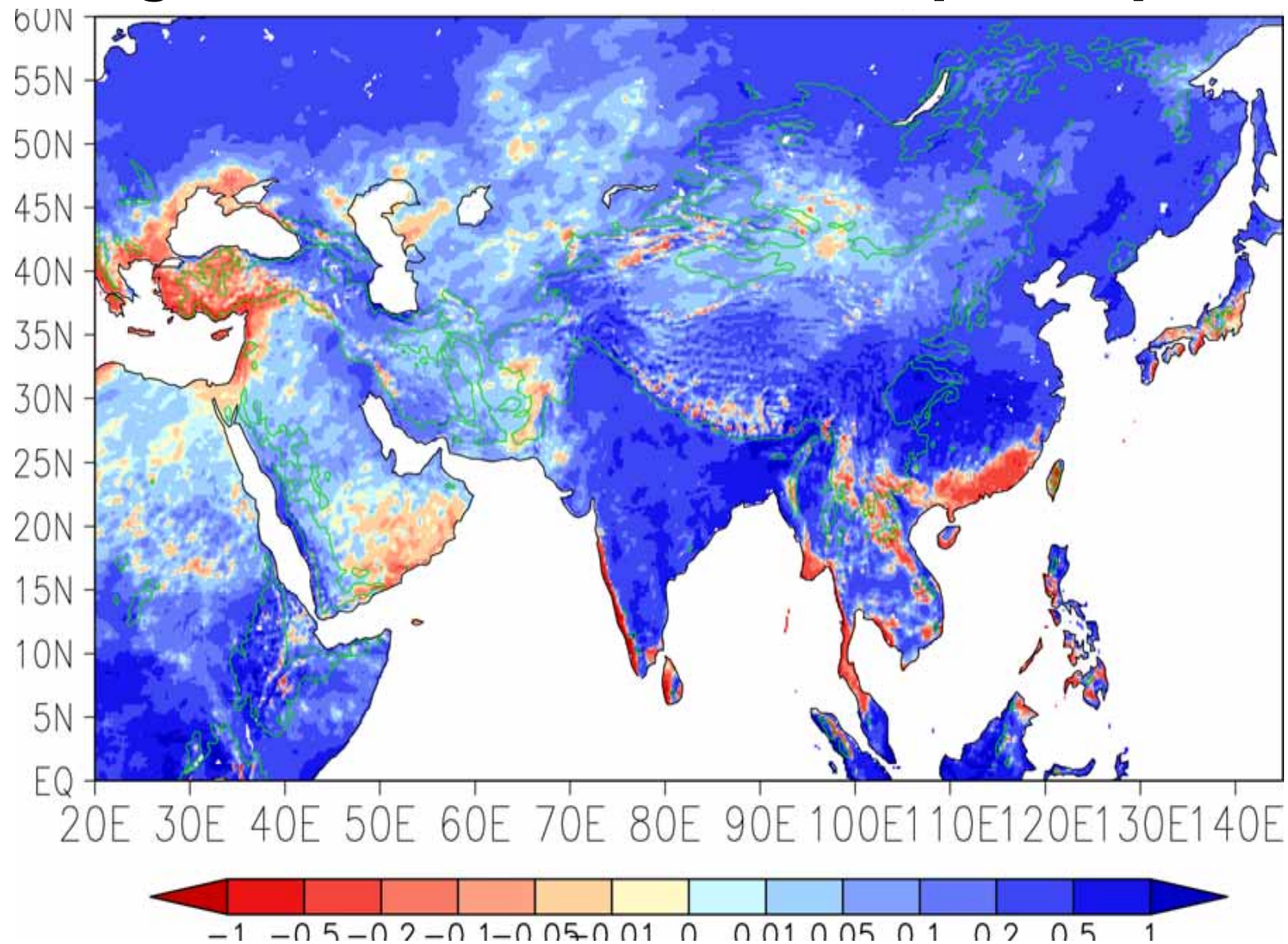


# JJA precipitation climatology



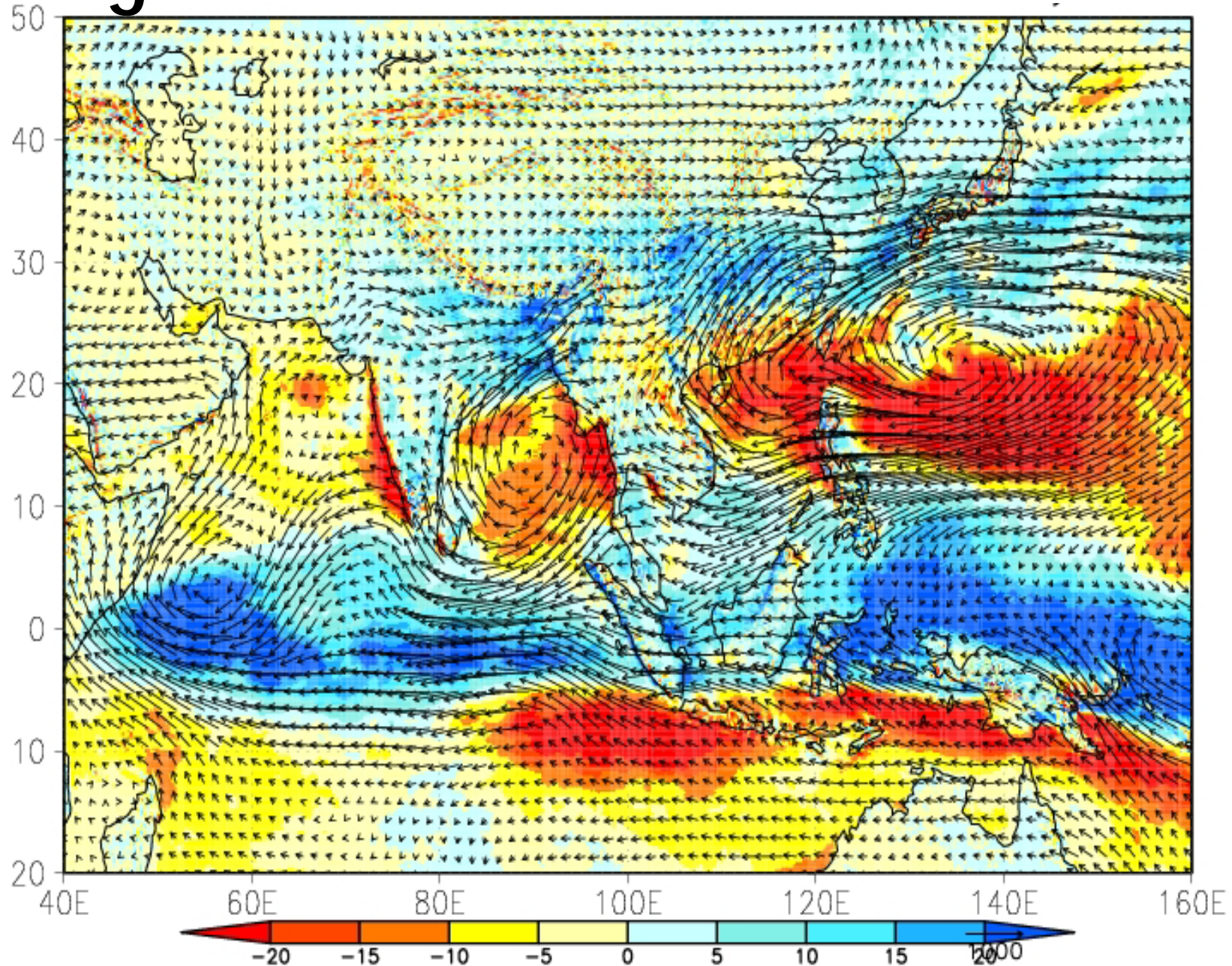


# Changes in annual mean precipitation

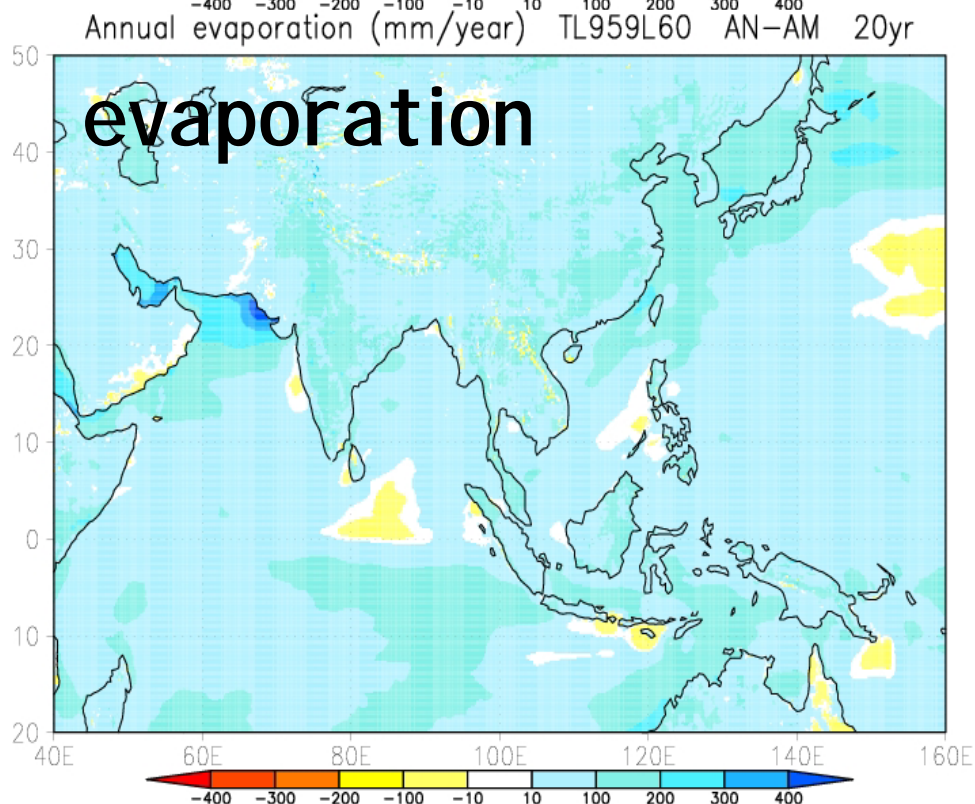
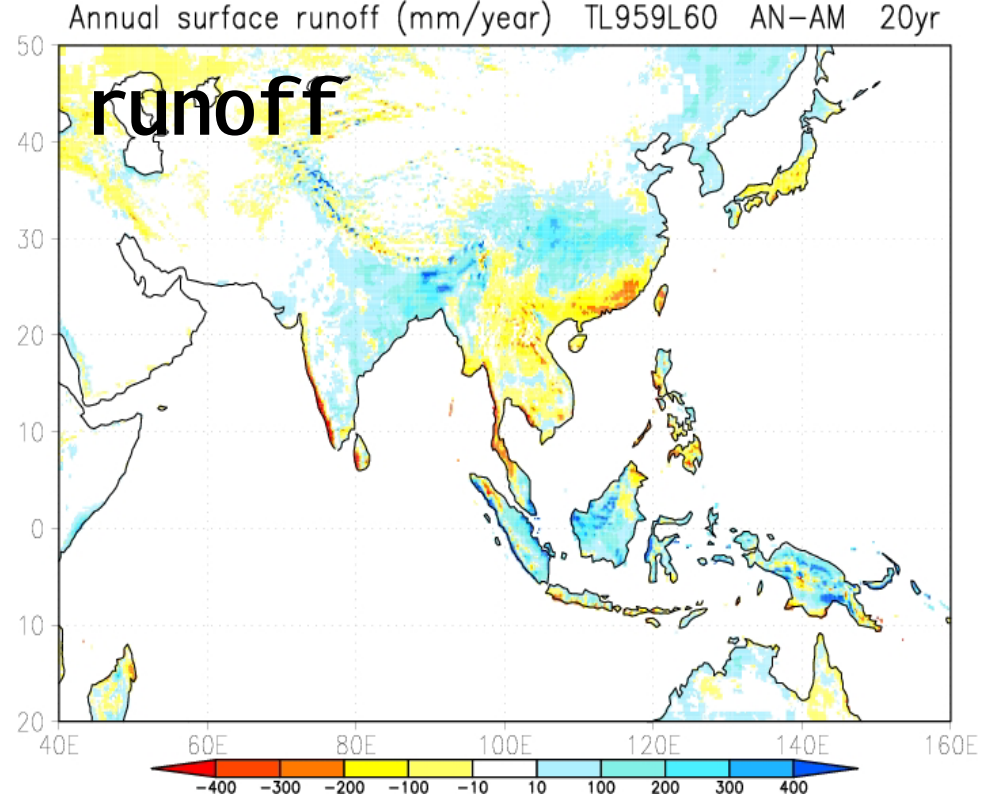
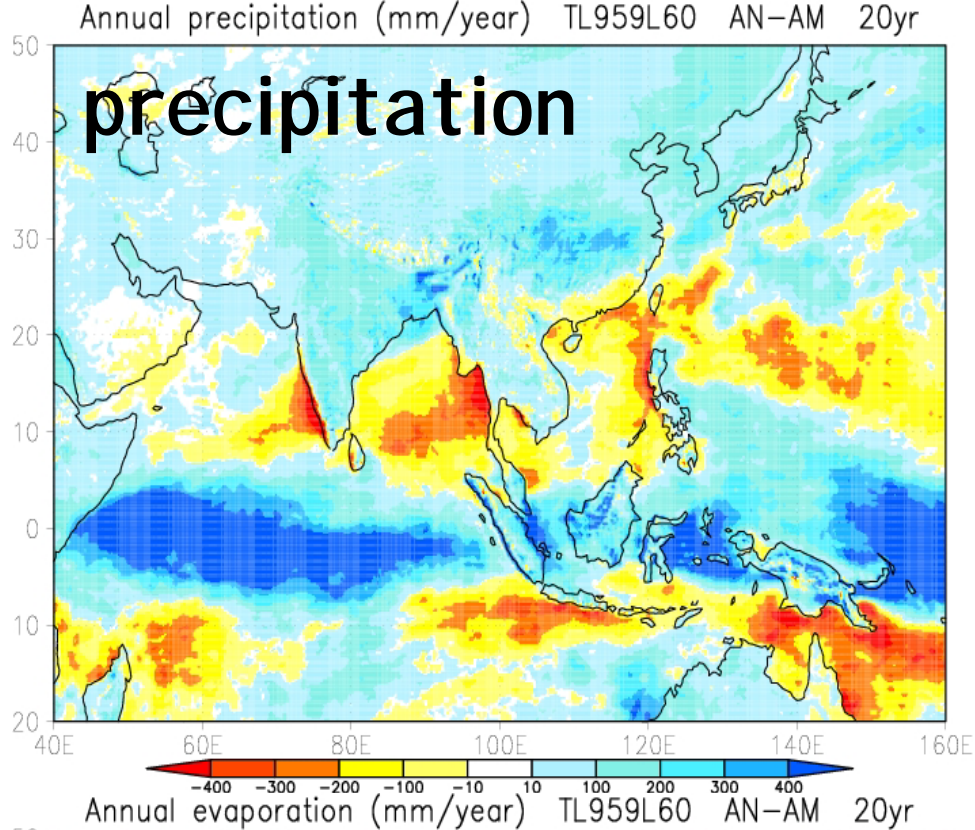


Overall increase in Asian precipitation, but there are areas with decreased precipitation, related to moisture flux changes ...

# Changes in JJA mean moisture flux

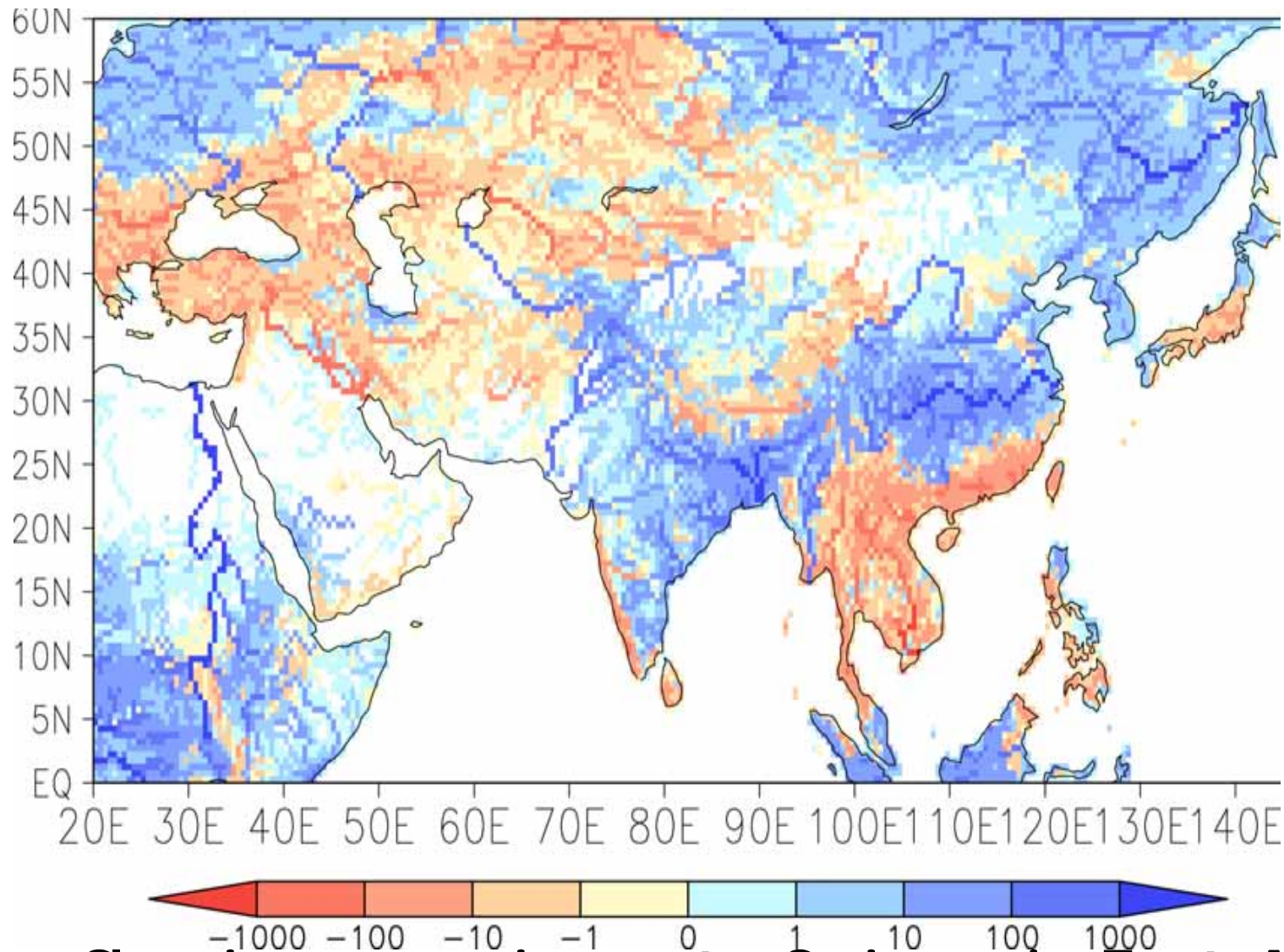


Over the oceans, large moisture flux divergence and convergence is found; mostly convergence over land



Due to general increase in evaporation, there are regions with decreased runoff though precipitation increases; an example is Thailand.

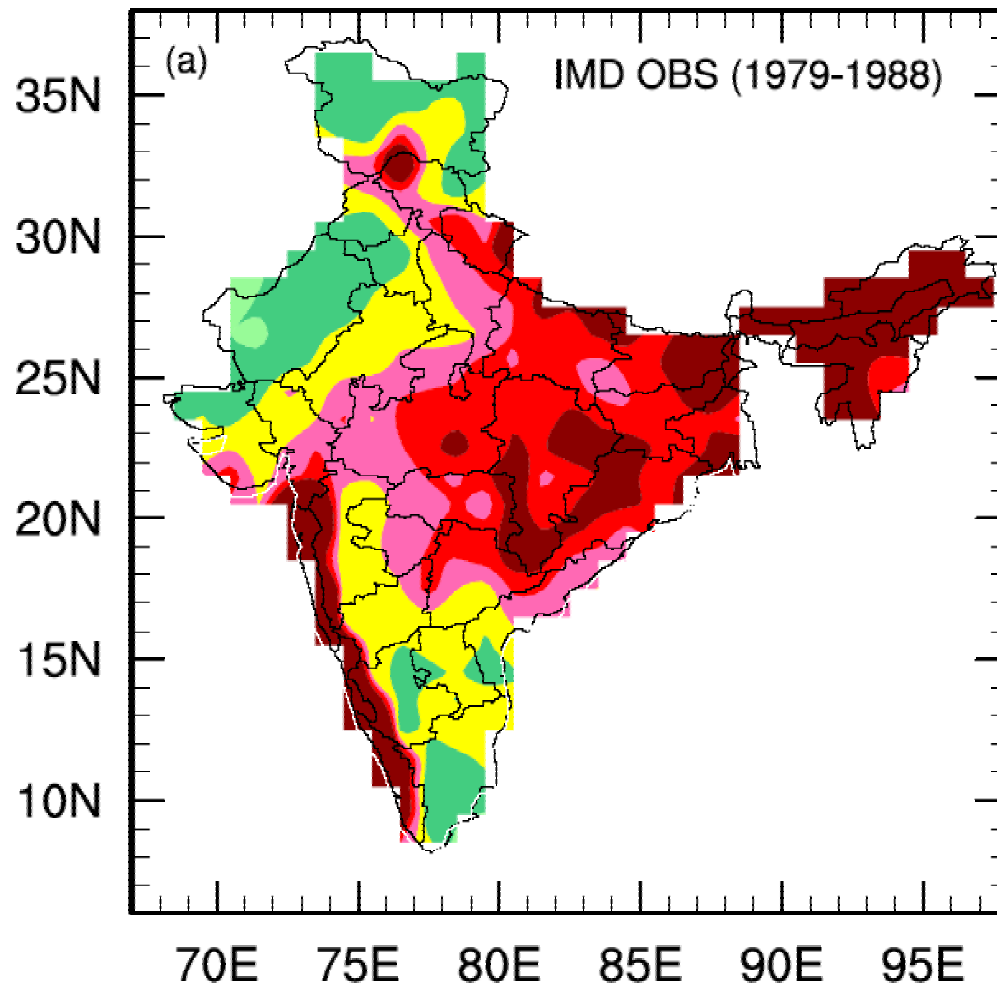
# Changes in annual streamflow



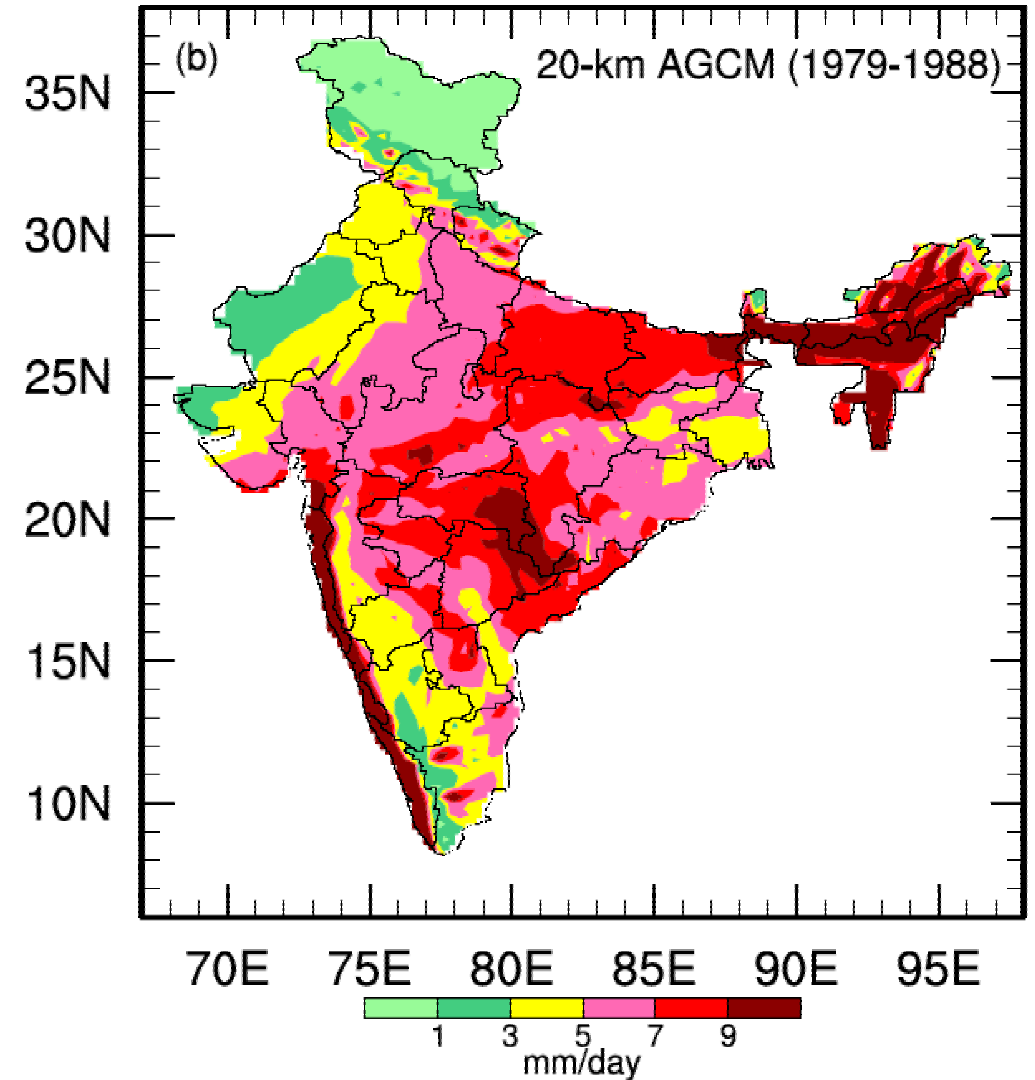
**Streamflow increases in most of rivers in East Asia and South Asia, but decreases in Southeast Asia, western Asia and the Mediterranean region**

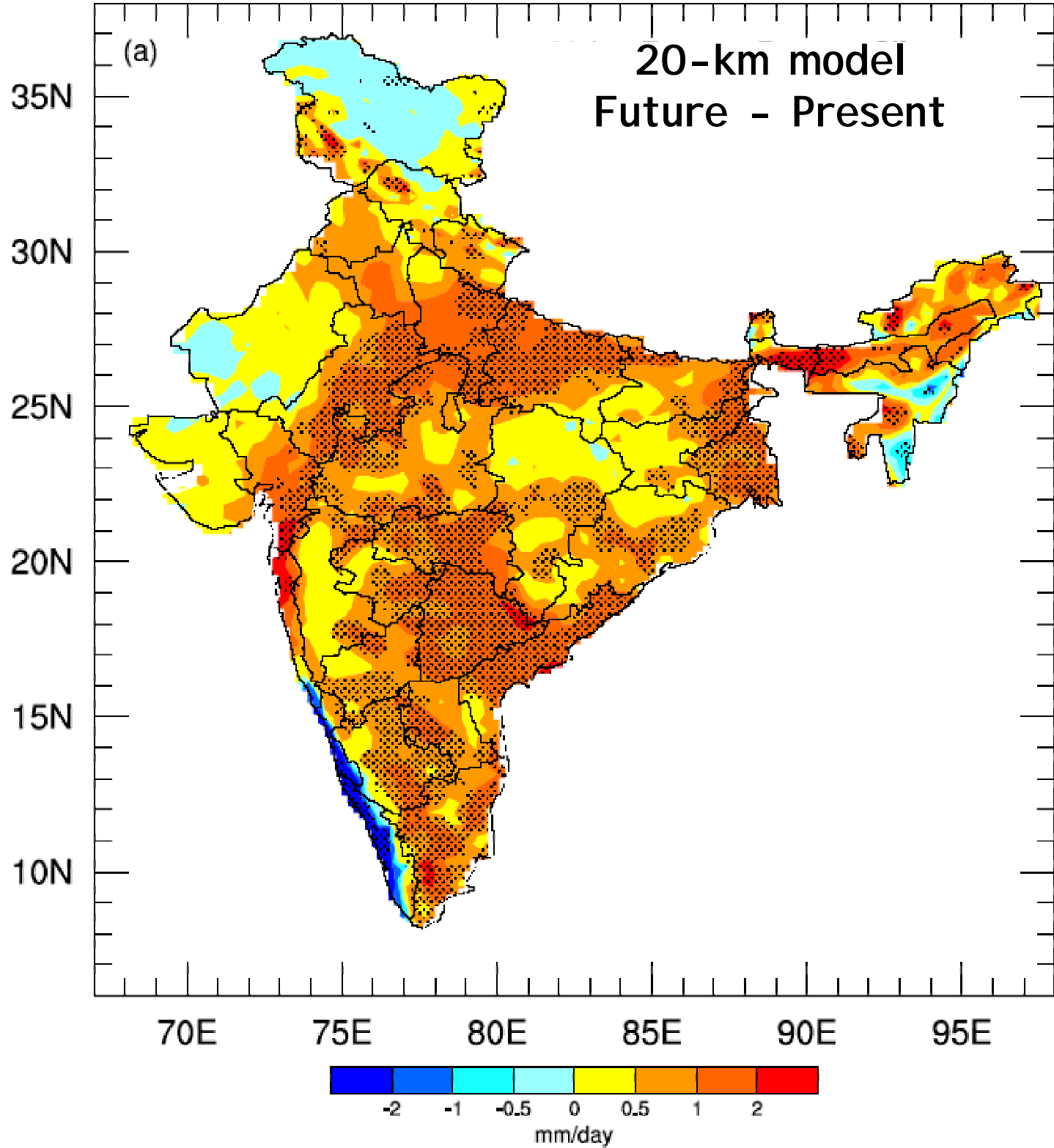
# Indian summer monsoon rainfall

IMD observation

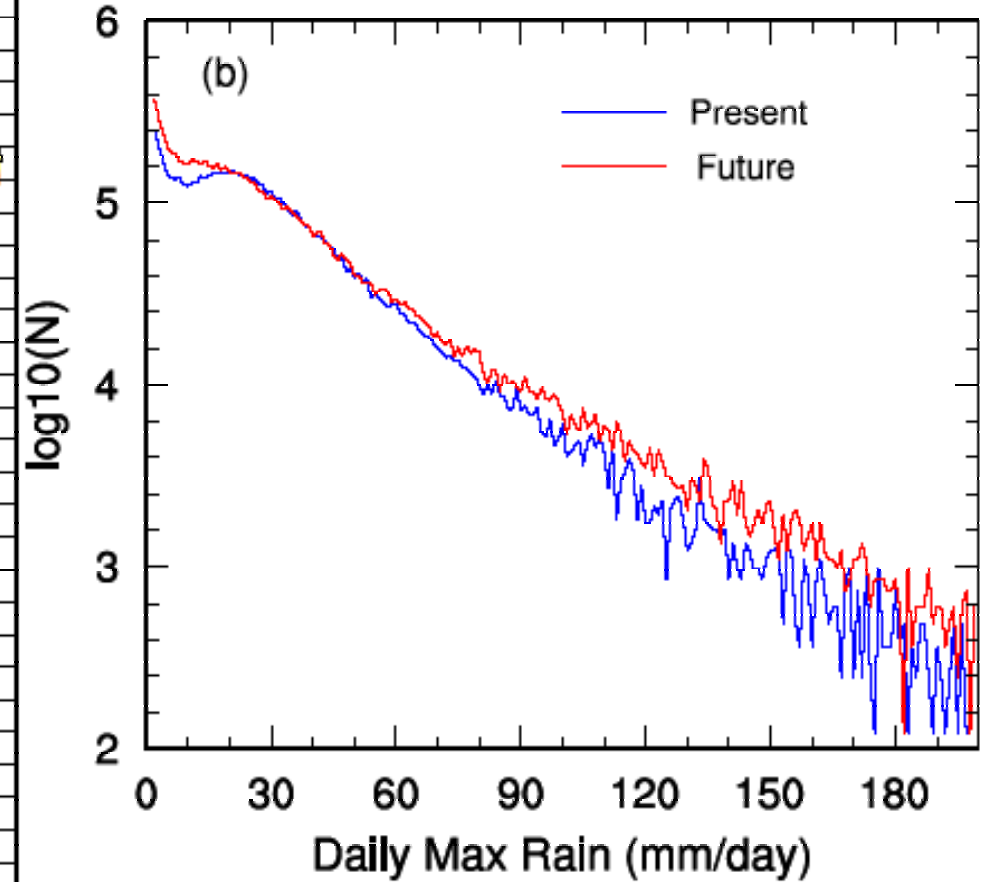
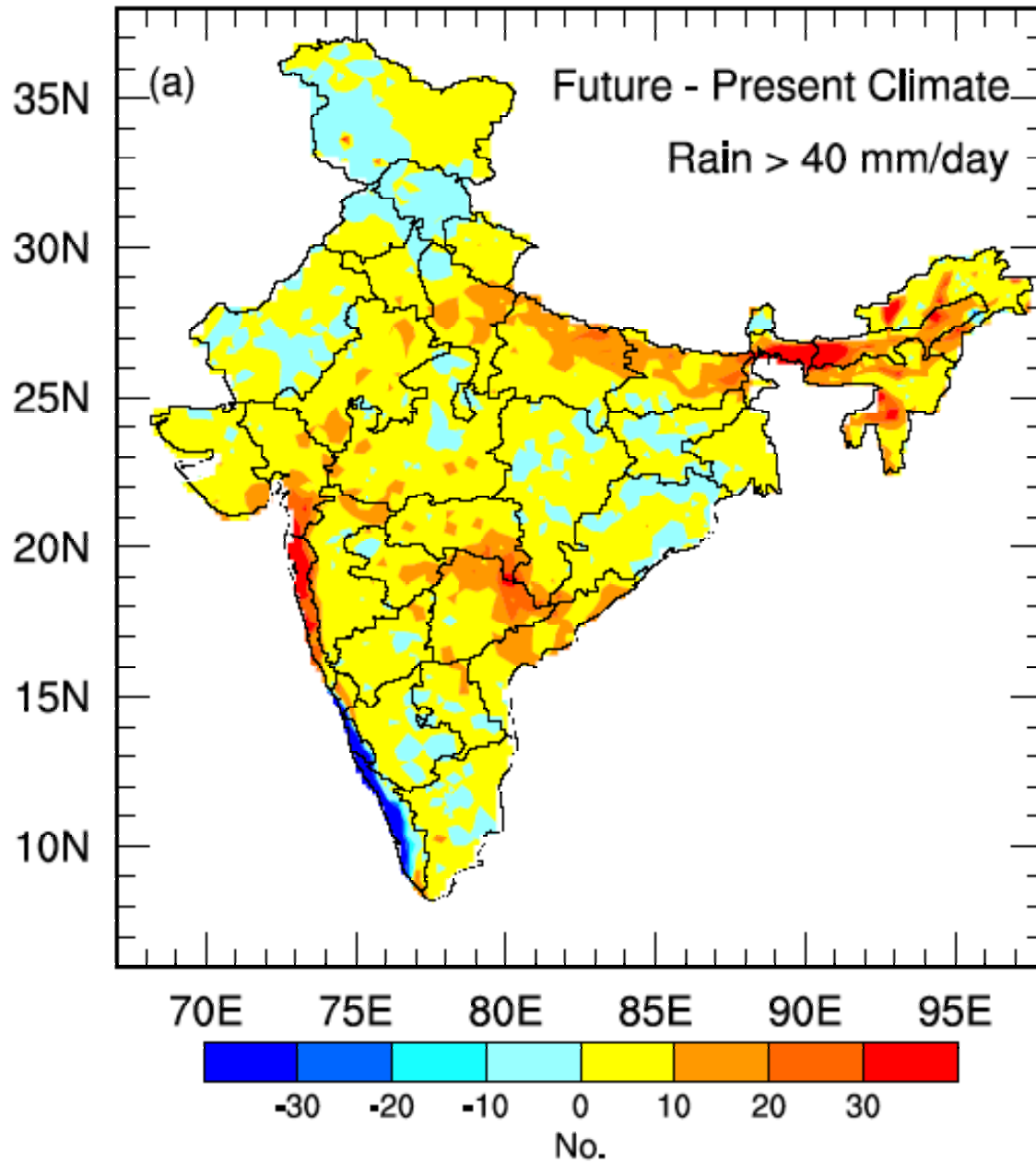


20-km model





# Heavy rainfall days

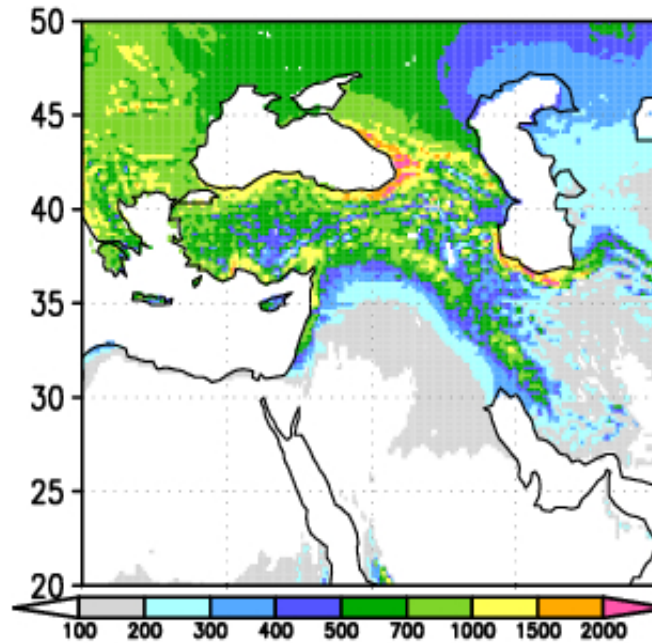
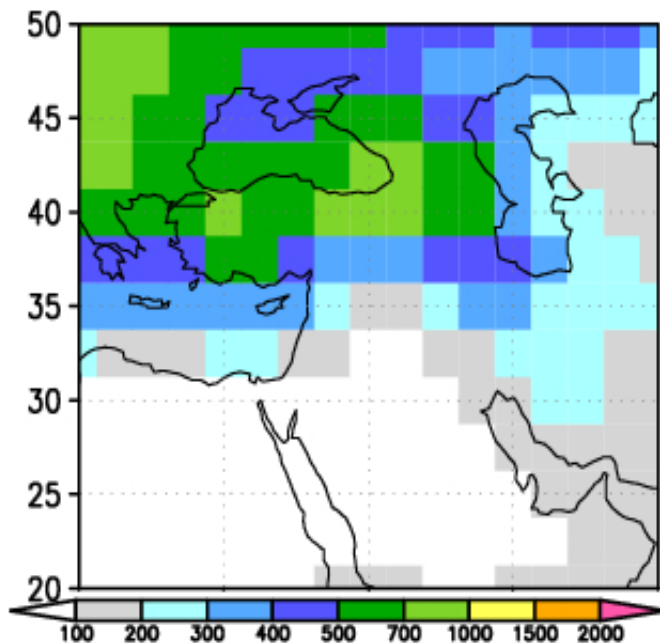


# Fertile Crescent: Annual Precipitation

CMI P3 models

20-km model

Present



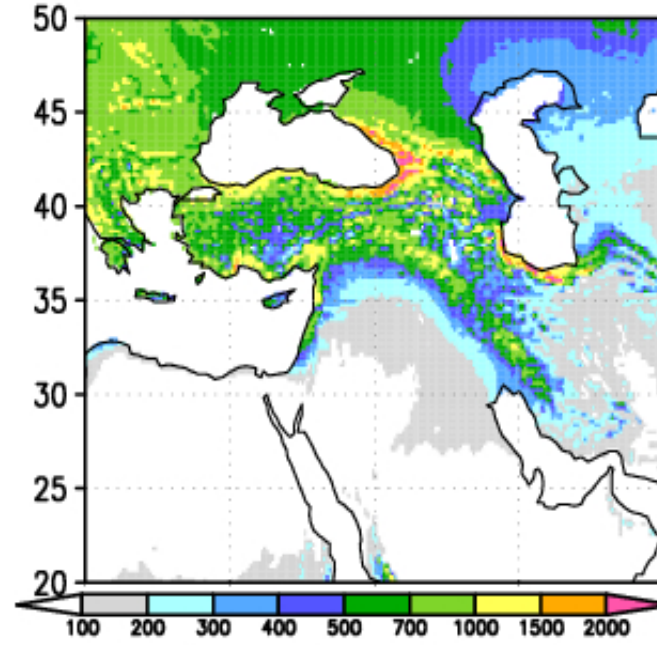
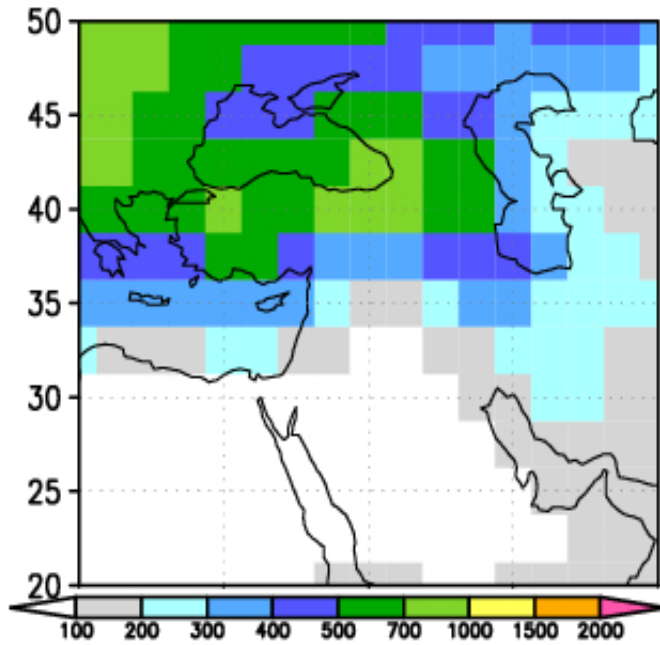


# Annual Precipitation

CMI P3 models

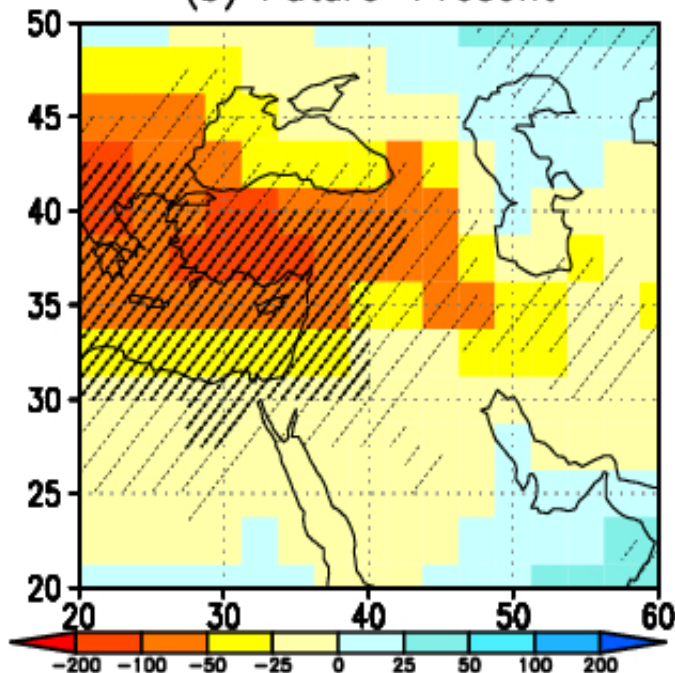
20-km model

Present

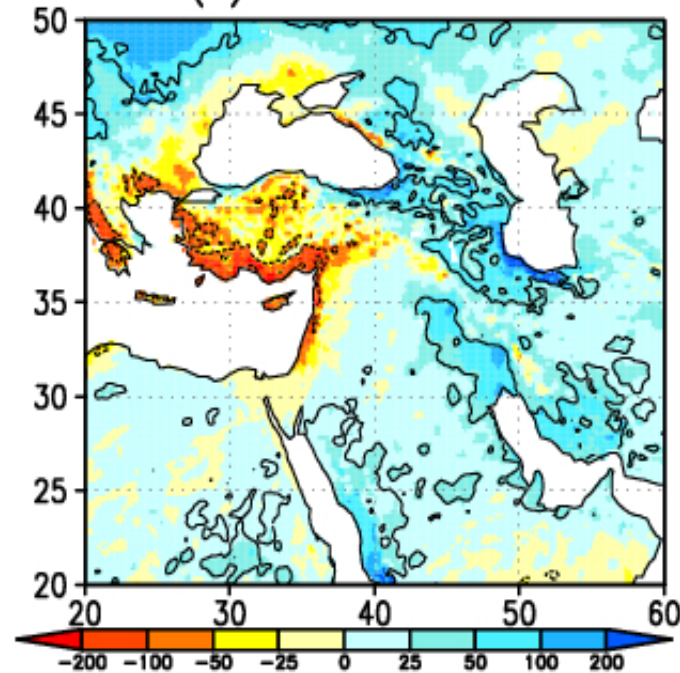


Future-Present

(b) Future-Present



(b) Future-Present



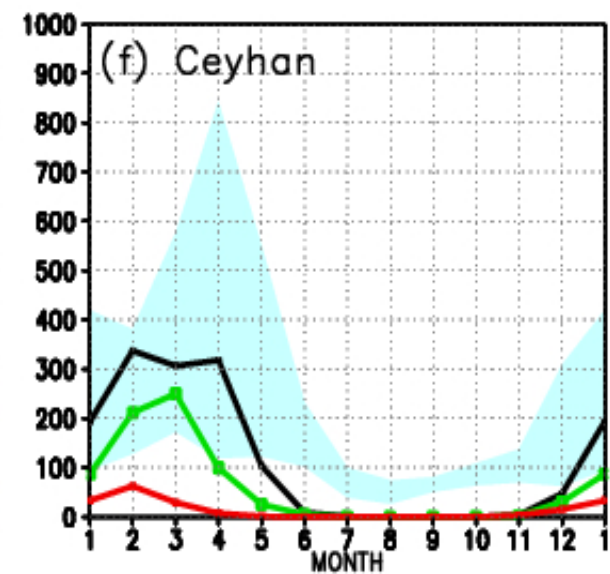
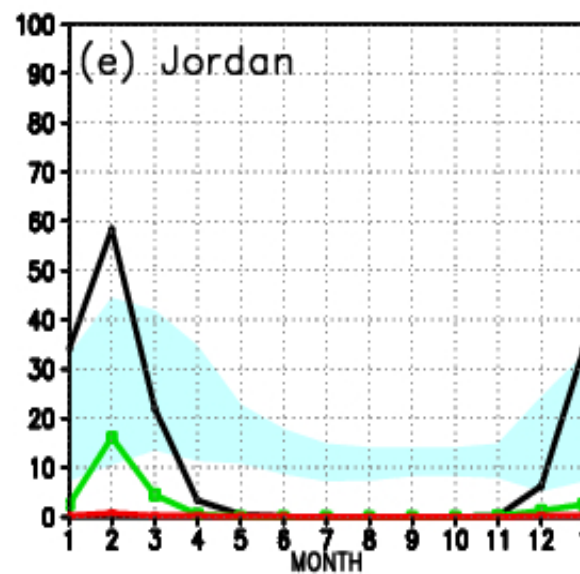
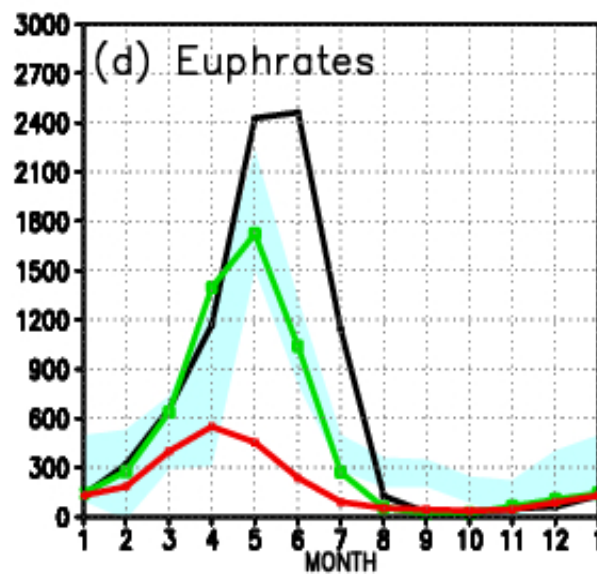
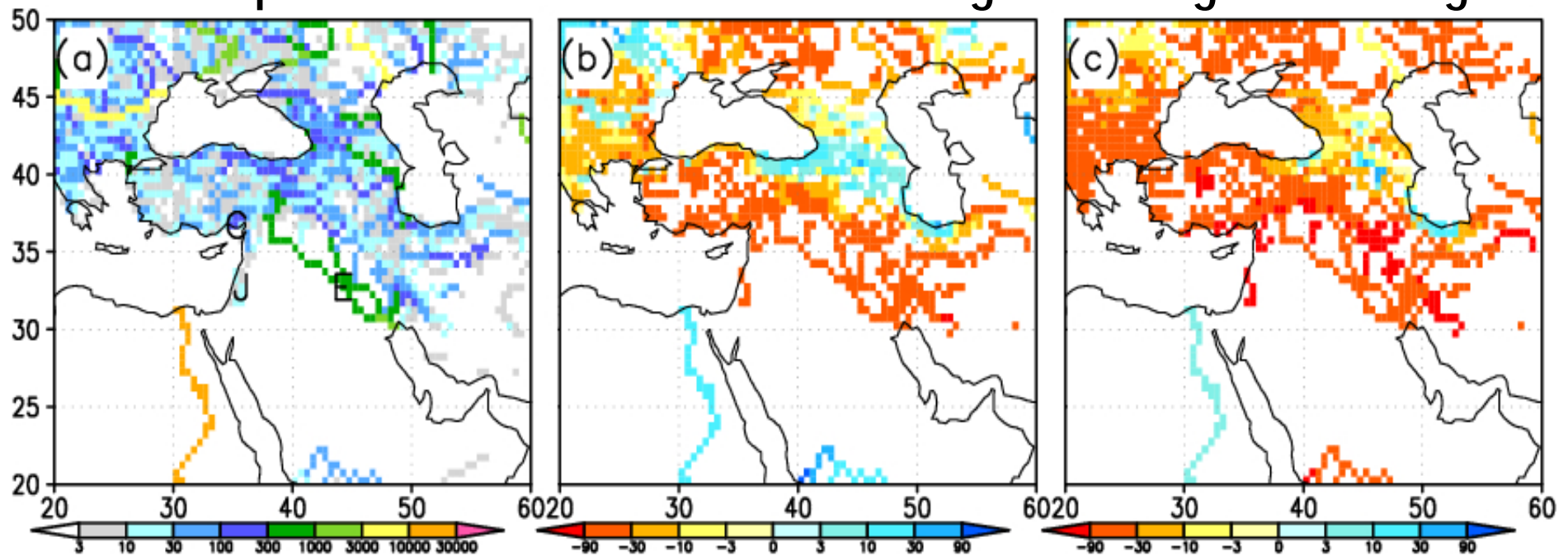
# Annual streamflow

2081-2100 changes

20km present

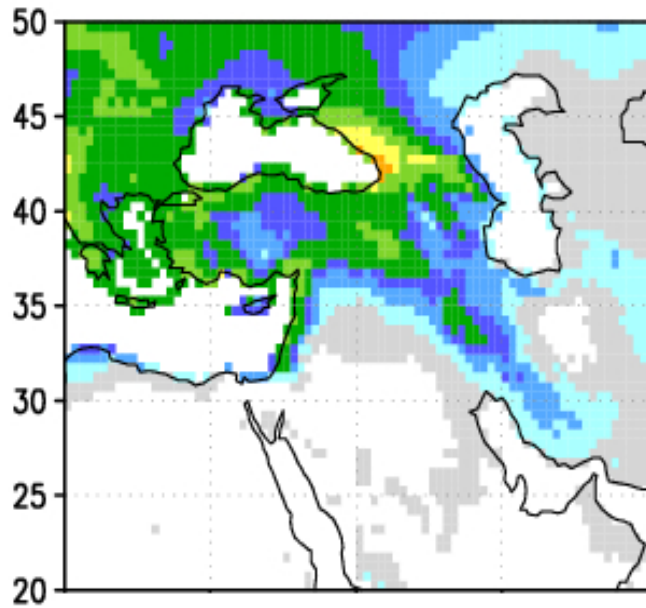
Medium Warming

High Warming

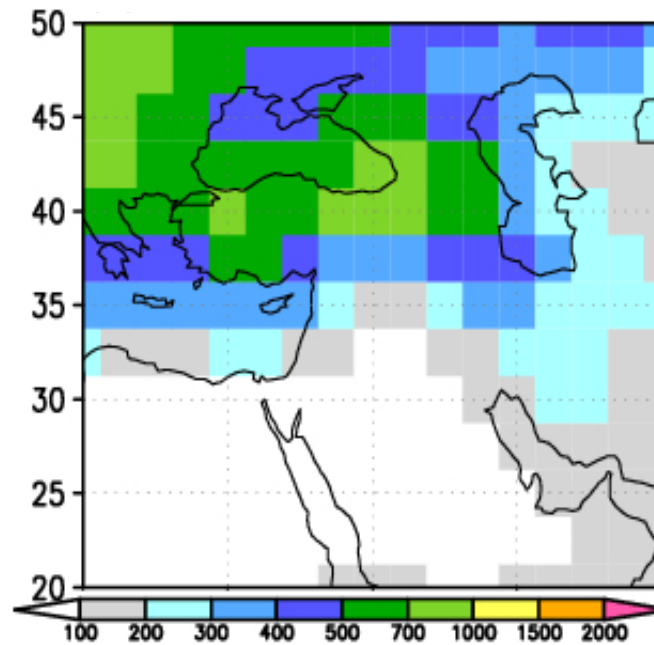


# Annual Precipitation

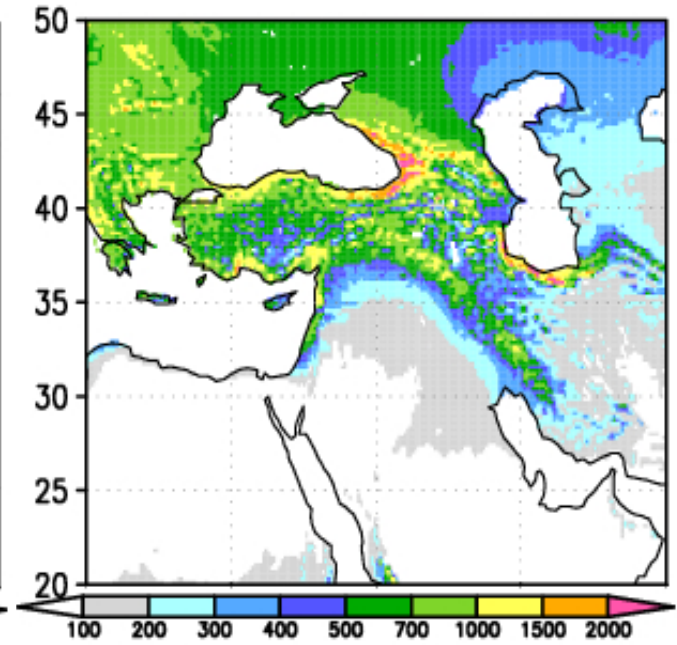
CRU (Observation)



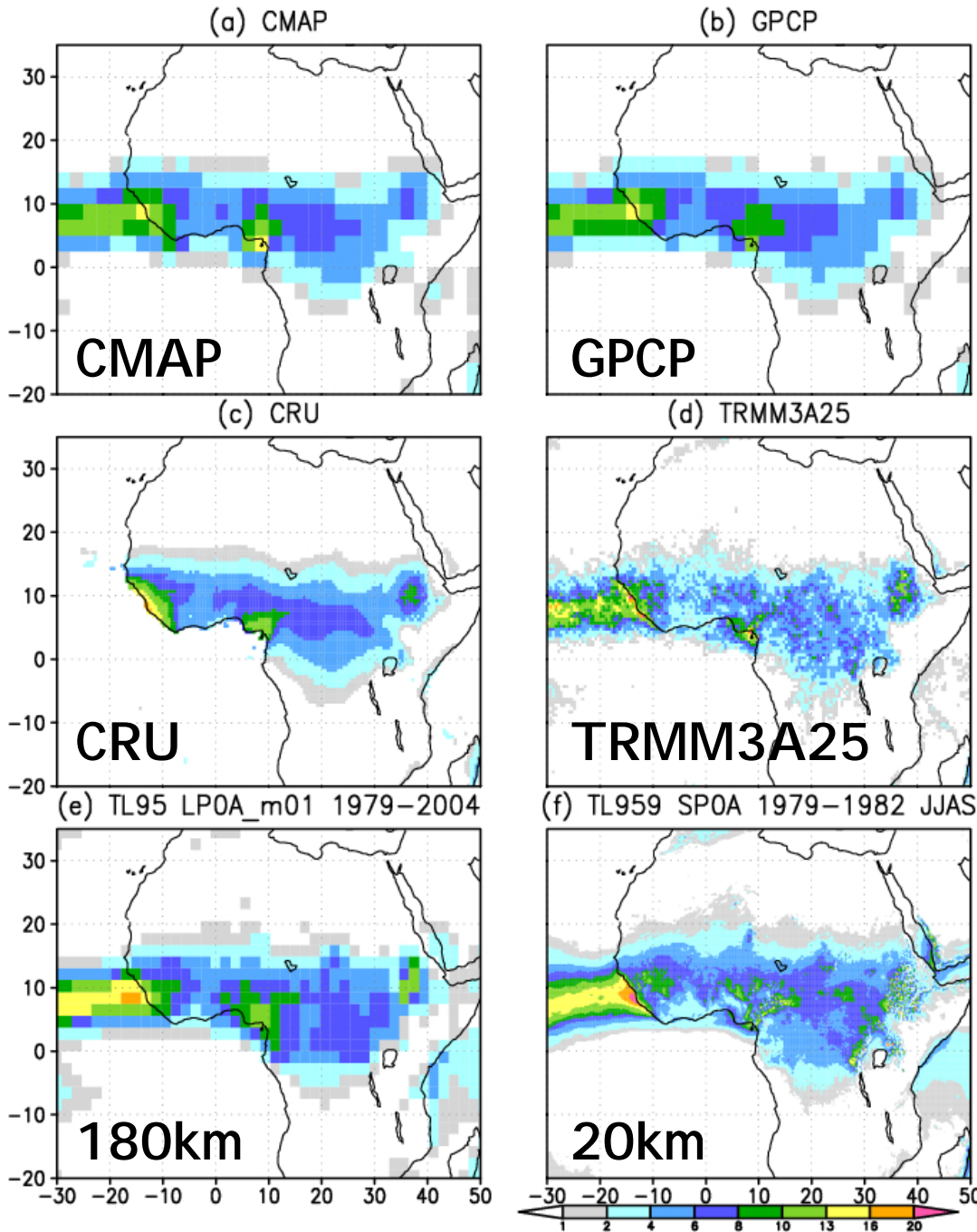
CMIP3 models



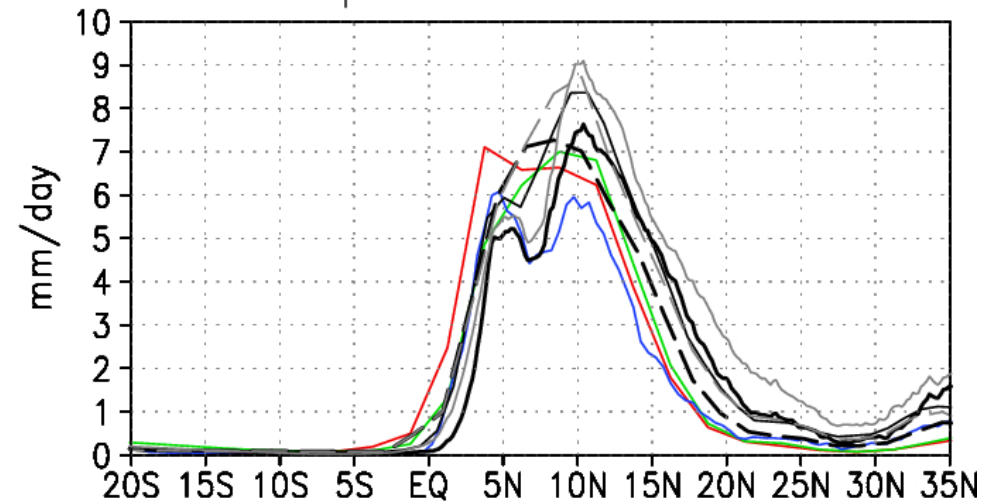
20-km model



# West African monsoon



JJAS Precip [10W-10E]

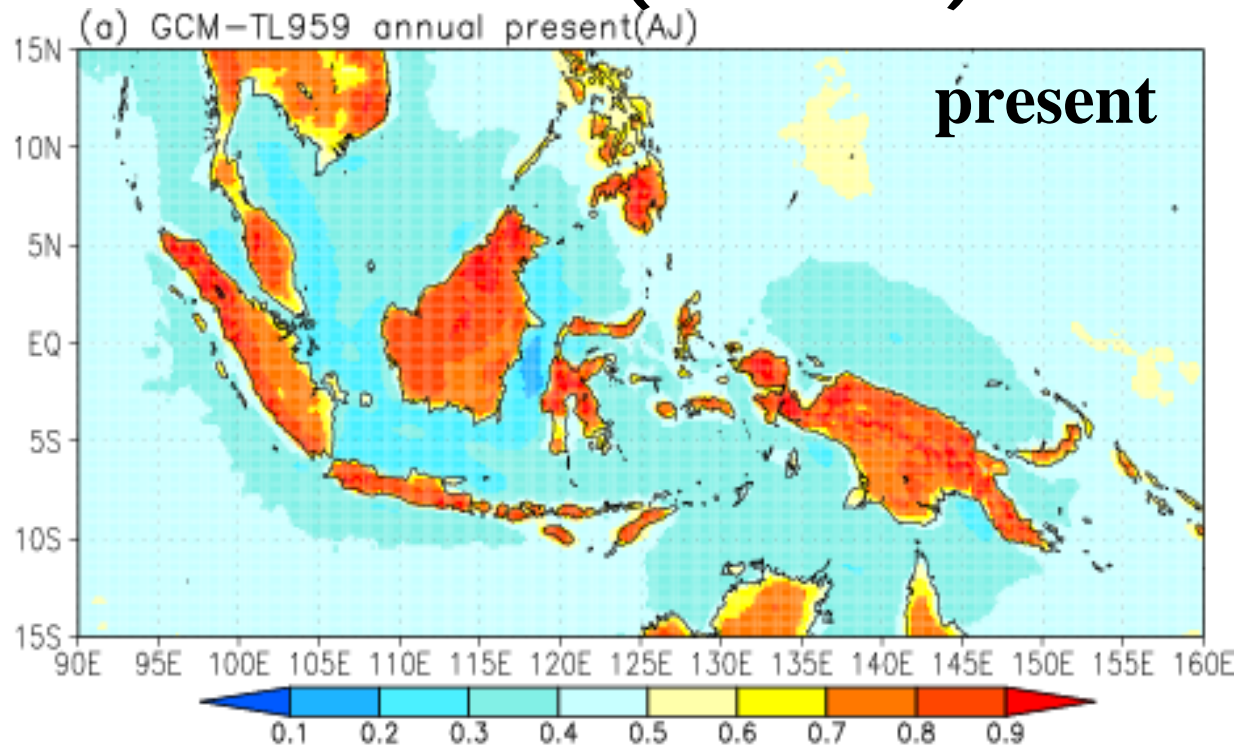


Black solid: 20km    Black dashed: 180km  
 Thin black solid: 120km  
 Red: CMAP    Green: GPCP    Blue: TRMM

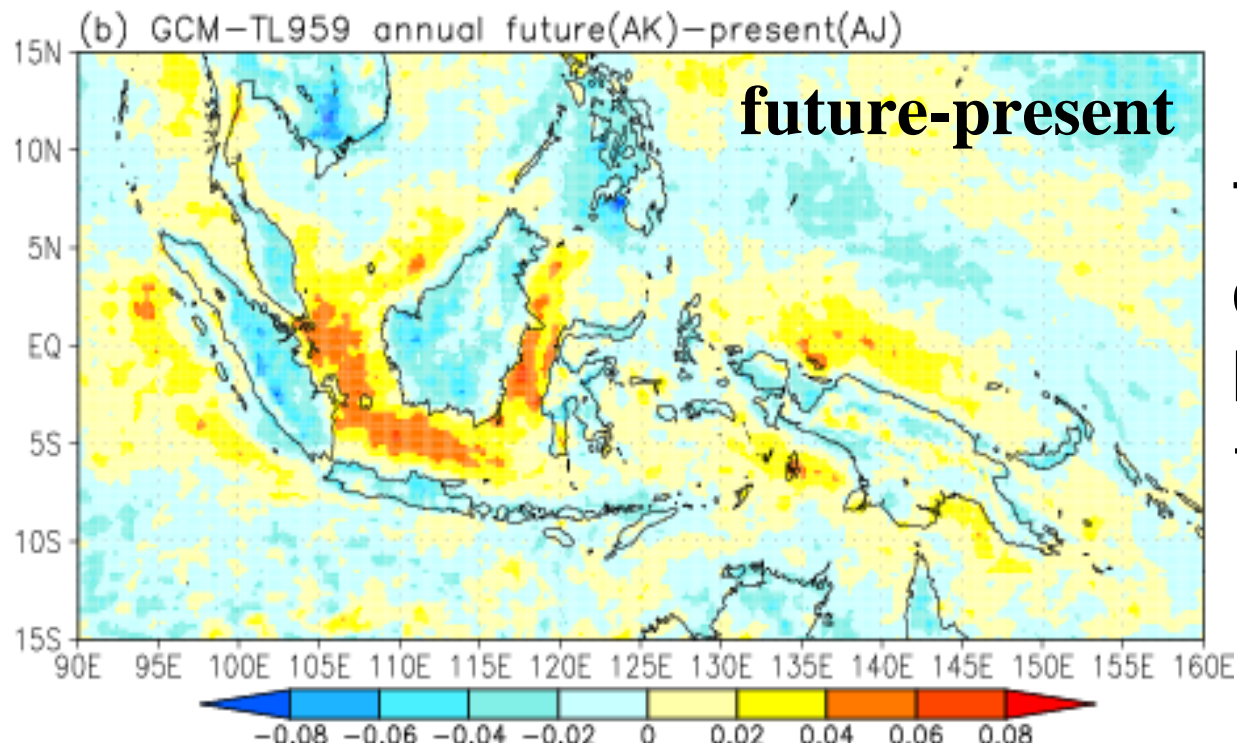
QuickTime<sup>®</sup> Ç²  
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AM rain  
vs  
PM rain

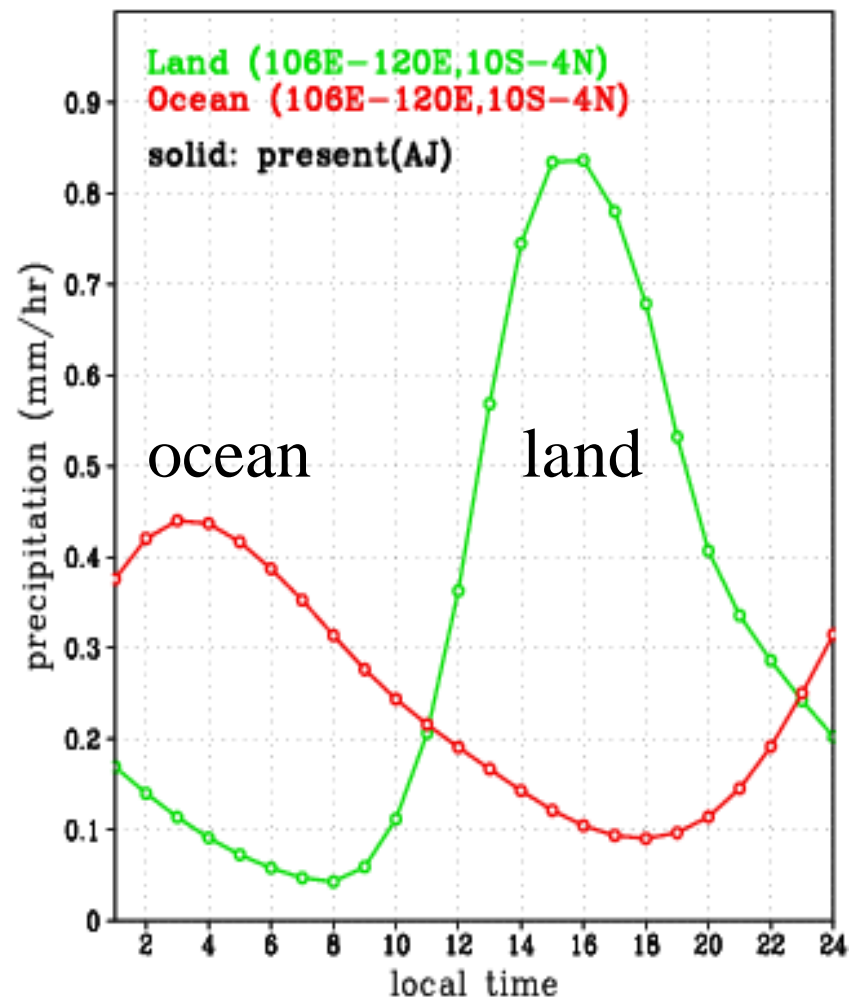
$PM / (AM+PM)$

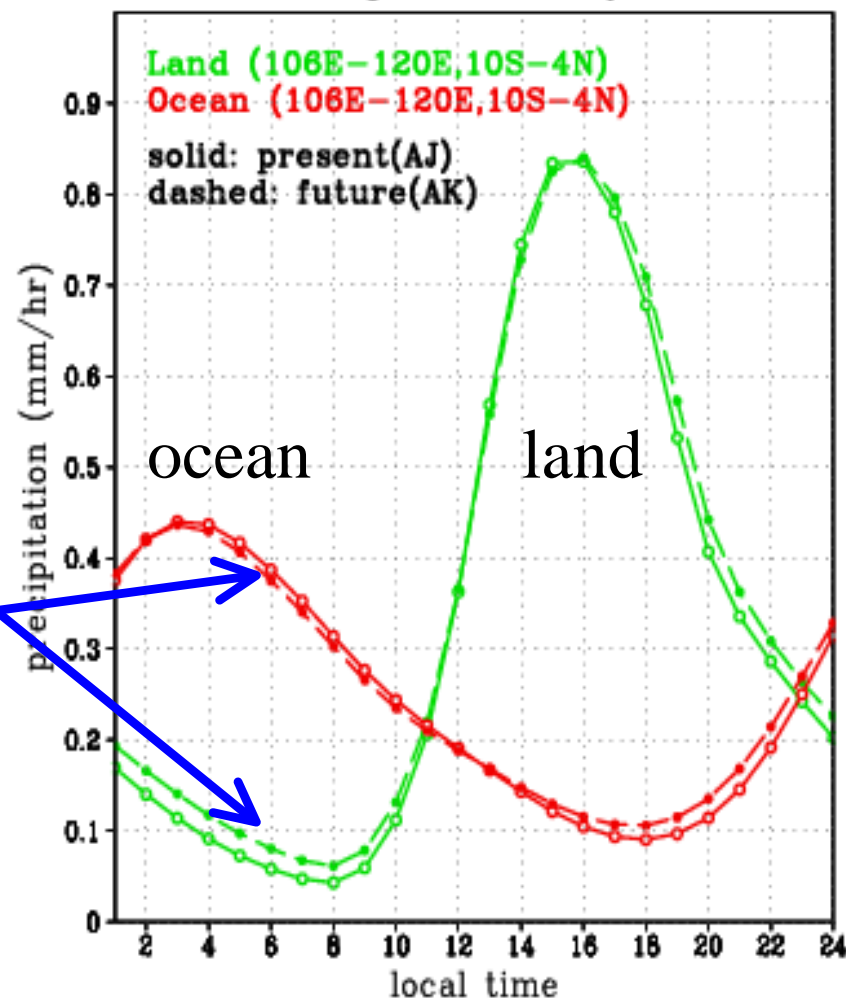


PM rain over  
land  
AM rain over  
the ocean



The ratio  
decreases over  
both land and  
the ocean

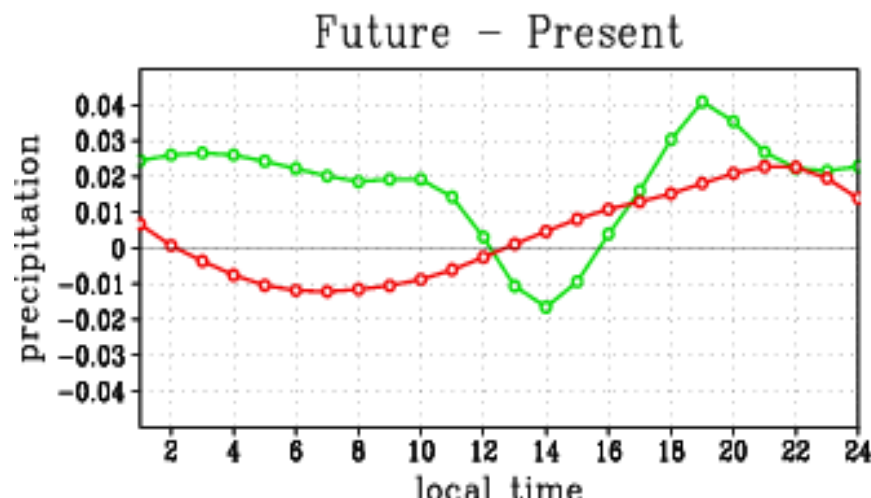




nighttime decrease  
over ocean and  
nighttime increase  
over land is noted

.

decrease in day-  
night precipitation  
contrast





# Summary

Resolution of climate models becomes finer; now we are using 60-km or even 20-km mesh global climate models.

But, for model validation, we still rely on conventional observed data, whose resolution is coarser than that of the model.

Therefore, reliable fine resolution verification data is highly needed.